

# **GRACE Follow-On**

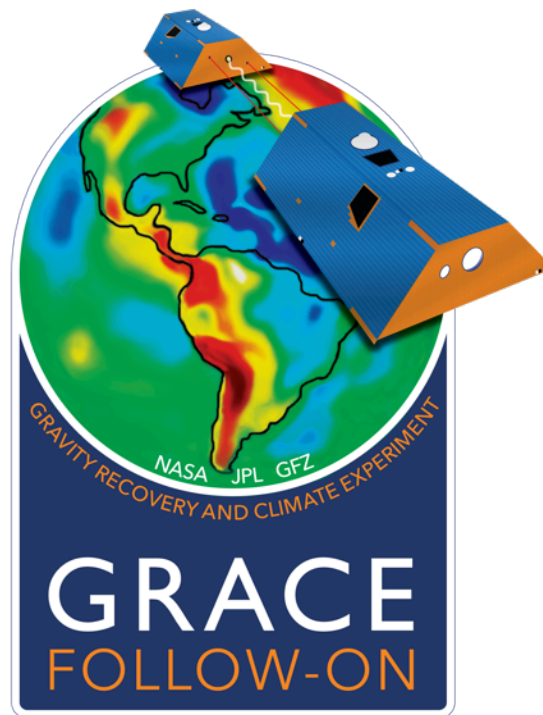
## **Gravity Recovery and Climate Experiment Follow-On**

### **CSR Level-2 Processing Standards Document**

#### **For Level-2 Product Release 06**

**CSR GRFO-19-01 (GRACE-FO D-103920)**

Himanshu Save  
Center for Space Research  
The University of Texas at Austin



Prepared by:

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Himanshu Save, UTCSR  
GRACE-FO Science Operations Manager

Contact Information:

Center for Space Research  
The University of Texas at Austin  
3925 W. Braker Lane, Suite 200  
Austin, Texas 78759-5321, USA  
Email: [grace@csr.utexas.edu](mailto:grace@csr.utexas.edu)

Reviewed by:

Srinivas Bettadpur, John Ries, and Peter Nagel  
UTCSR, Austin

Model Development, Coding & Verification by: CSR Level-2 Team

**Document Change Record**

<b>Issue</b>	<b>Date</b>	<b>Pages</b>	<b>Change Description</b>
01.0	June 6, 2019	All	Initial Version
01.1	June 11, 2019	13-14	Revision to add LRA offset model

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## I INTRODUCTION

### I. 1 PURPOSE OF THE DOCUMENT

This document serves as a record of the processing standards, models & parameters adopted for the creation of the Level-2 gravity field data products by the GRACE Follow-On (GRACE-FO) Science Data System component at The University of Texas Center for Space Research (UTCSR). This document is issued once for every release of Level-2 data products generated by UTCSR. That release number is included in the title of this document. The release number refers to the field *rr* in the generic Level-2 product name (see *Product Specification Document* or *Level-2 Gravity Field Product User Handbook*)

*PID-2\_YYYYDOY-YYYYDOY\_ddd\_UTCSR\_mmmm\_rrvv*

Where *PID* is 3-character product identification mnemonic, *YYYYDOY-YYYYDOY* specifies the date range (in year and day-of-year format) of the data used in creating this product, *ddd* specifies the gravity mission, *mmm* is a 4-character mnemonic used to identify the characteristics of the gravity solution, *rrvv* specifies 2-digit release number and 2-digit version number.

This document may be used in conjunction with:

1. GRACE Follow-On Level-1 User Handbook (JPL D-56935)
2. GRACE Follow-On Level-2 Gravity Field Product User Handbook (JPL D-103922)
3. GRACE Follow-On JPL L-2 Processing Standards Document (JPL D-103921)
4. GRACE Follow-On GFZ L-2 Processing Standards Document (JPL D-103919)
5. GRACE AOD1B Product Description Document for AOD1B Release 06 (327-750)

### I. 2 DOCUMENT CHANGE HISTORY

This document is the first release.

### I. 3 OVERVIEW OF DATA PROCESSING

This section contains a brief overview of the data processing done to obtain the Level-2 products in this release.

The gravity field estimates were made using the conventional dynamic, linear least squares adjustment for the orbit and gravity field from an optimally weighted combination of the GPS & K-Band tracking data collected by the GRACE-FO satellites. Some specifics follow in the next table.

<b>Processing Institution</b>	<b>University of Texas Center for Space Research</b>	
<b>Software Used</b>		
Orbit Software	MSODP	Version 2019.1

Linear Solver Software	AESoP	Version 20120322_v001
<b>GRACE-FO Data Products Used</b>		
<i>Product ID &amp; Release</i>	<i>Data Rate</i>	<i>Remarks</i>
ACT1B (RL=04)	1 second	Used in the numerical integrator
SCA1B (RL=04)	1 second	For observation models & transforming body-fixed accelerations
KBR1B (RL=04)	5-second Range Rate only	
GPS1B (RL=04)	2-minute Double Differences between one GRACE-FO satellite, two GPS satellites & one ground station	IGS core-station network. IGS14 Orbits produced by JPL used for GPS satellites – and held fixed during analysis
AOD1B (RL=06)	Used as part of background gravity acceleration models	
<b>Other Notes on Methodology</b>		
<p>Solution obtained as a weighted combination of GPS double differences for each satellite and inter-satellite K-Band Range-Rate – using one-day dynamic arcs over the prescribed data span. GPS data weight was limited to 2 cm for each double difference observation. K-Band range-rate was allowed optimal weighting.</p>		
<p>The project operational product is the outcome of the unconstrained linearized least-squares estimation where gravity parameters were estimated independently from the other orbit related parameters estimated during an earlier iteration. The regularized version of the product is experimental, and the regularization is not described in this document.</p>		

## II ORBIT DYNAMICS MODELS

### II.1 EQUATIONS OF MOTION

The equations of motion for both GRACE-FO satellites are identical in mathematical form. In the remainder of this chapter, the equations will be provided for a single Earth orbiting satellite, with the understanding that the same equations apply to both GRACE-FO satellites. Where appropriate, the parameters or conditions unique to each satellite will be specified.

In the inertial frame

$$\ddot{\vec{r}} = \vec{f}_g + \vec{f}_{ng} + \vec{f}_{emp}$$

where the subscript “g” denotes gravitational accelerations; “ng” denotes the acceleration due to the non-gravitational or skin forces; and “emp” denotes certain empirically modeled forces designed to overcome deficiencies remaining in the force models.

#### II.1.1 Time Systems

The independent variable in the equations of motion is the time system TDT (Terrestrial Dynamical Time). The relationship of this abstract, uniform time scale to other time systems is well defined. The table below shows the relationship between various time systems and the contexts in which they are used.

System	Relations	Notes	Standards
TAI	Fundamental time system	International Atomic Time	n/a
UTC	UTC = TAI – n1 (Time-tag for saving intermediate products)	n1 are the Leap Seconds	Tables from USNO
TDT	TDT = TAI + 32.184s	This is the independent variable for numerical integration.	IAU 1976 Recommendation (equivalent to using TT in <i>IERS-2010</i> )
GPS	GPS = TAI – 19s (time-tag of GRACE-FO observations)	Relationship between GPS & TDT is fixed at 19 seconds	Time-tags in sec since 1200 Jan 01, 2000 GPS Time.

#### II.1.2 Reference Frames

The fundamental reference frame for the mathematical model is the non-rotating, free-falling (inertial) reference frame, with the origin defined as the center of mass of the Earth system. The Inertial and Earth-fixed reference frames, and their relative orientations and associated standards are further described in the chapter on Earth Kinematics.

## II. 2 GRAVITATIONAL ACCELERATIONS

The gravitational accelerations are the sum of direct planetary perturbations and the geopotential perturbations. The vector of direct planetary perturbations is evaluated using the planetary ephemerides. All geopotential accelerations are represented using a spherical harmonic expansion with time-variable coefficients, to a specified maximum degree and order. The accelerations are computed by evaluating the Earth-fixed gradient of the geopotential, which are then rotated (after summation with the non-gravitational accelerations) to inertial frame for the integration of equations of motion. In general,

$$\vec{f}_g = {}_{3 \times 3} M_{ef}^{in}(P, N, R) \vec{f}_g^{ef}$$

The 3x3 rotation matrix M, which depends on Earth orientation is described in the chapter on Earth Kinematics.

Contributions to the spherical harmonic coefficients of the geopotential, and the associated implementation & standards are now compiled. The geopotential at an exterior field point, at time t, is expressed as

$$U_s(r, \varphi, \lambda; t) = \frac{GM_e}{r} + \frac{GM_e}{r} \sum_{l=2}^{N_{\max}} \left( \frac{a_e}{r} \right)^l \sum_{m=0}^l \bar{P}_{lm}(\sin \varphi) [\bar{C}_{lm}(t) \cos m\lambda + \bar{S}_{lm}(t) \sin m\lambda]$$

where r is the geocentric radius, and  $(\varphi, \lambda)$  are geographic latitude and longitude, respectively, of the field point.

The suite gravitational models used for propagation of the equations of motion of the satellites are called the Background Gravity Models. This concept, and its relation to GRACE-FO estimates, is described further in the *Level-2 User Handbook*. The details of the background gravity model are provided here.

Hereafter, the document *IERS Conventions (2003)* is abbreviated as *IERS-2003*, and the *IERS Conventions (2010)* as *IERS-2010*.

### II.2.1 Mean Geopotential & Secular Changes

Parameter	Value	Remarks
$GM_e$	3.986004415E+14	<i>IERS-2010 Standards (value is consistent with using TDT or TT as the time argument)</i>
$a_e$	6378136.3 m	
$N_{\max}$	Complete to degree and order 360	<b>GGM05C</b> is background static model.
Secular Change	N/A	Not modeled in background.

Note 1: The normalization conventions are as defined in *IERS-2010*, Eqs 6.2-6.3.

Note 2: The implementation of computation of spherical harmonics and its derivatives is



as described in (Lundberg and Schutz, 1988).

Note 3: The degree-1 terms are exactly zero in the geopotential model.

Note 4: The mean field **GGM05C** is a mean gravity field model created from a combination of approximately 10 years of RL05 products spanning from March 2003 through May 2013. The GRACE data were combined with GOCE gradiometer data from November 2009 to October 2013. In addition, harmonic coefficients extracted from the DTU13 gravity anomaly dataset extended GGM05C to degree and order 360, as described in *Ries et al. (2016)*. The model coefficients are available from GRACE data archives, ICGEM, and from <ftp://ftp.csr.utexas.edu/pub/grace/GGM05>

### II.2.2 Solid Earth Tides

Solid Earth tidal contributions to the geopotential are computed for the an-elastic Earth model, as specified in Section 6.2, *IERS Conventions (2010)*. Corrections to specific spherical harmonic coefficients are computed and added to the mean field coefficients.

Model	Notes	
Planetary Ephemerides	DE-430	
Frequency Independent Terms (an-elastic Earth)	Degree 2 & 3 – expression in Eq 6.6, <i>IERS-2010</i> .	Parameter values from Table 6.3
	Ellipticity contributions from Degree 2 tides to Degree 4 terms	As per Eq. 6.7, <i>IERS-2010</i>
Frequency Dependent Terms	Corrections to all degree-2 terms	As per Tables 6.5, <i>IERS-2010</i>
Permanent Tide in $\bar{C}_{20}$ (zero-tide system)	4.201E-9	Subtracted from total contributions as calculated above (implicitly included in value of the mean C20)

### II.2.3 Ocean Tides

The ocean tidal contributions to the geopotential are computed as specified in the prefatory material for Section 6.3, *IERS-2010*. Corrections to specific spherical harmonic coefficients of arbitrary (selectable) degree and order are computed and added to the mean field coefficients. The background ocean tide models are detailed in the table below, and the models and the methods of interpolation to minor tidal constituents are different from the model and methods specified in *IERS-2010*.

Model	Description	Notes
Tidal Arguments & Amplitudes/Phases	<i>Doodson (1921)</i> <i>Cartwright &amp; Tayler (1971)</i>	
Diurnal/Semi-Diurnal Bands	Harmonics of model <b>GOT4.8</b> ( <i>Ray 2012, pers. comm.</i> ) to degree 180 (See Note-1)	Extended to all minor constituents by fitting
	Periods > Monthly: Self-consistent equilibrium model	

Long-Period Band	(Ray 2005 pers. comm. based on Ray & Cartwright 1994)	admittances to the provided estimates for the major tides.
	Mm and Mf: Egbert & Ray 2003	
	Mtm and Msm: FES2004 (Lefevre 2005)	
Note-1: GOT4.8 differs from GOT4.7 in only the harmonics of the S2 tide.		

The implementation is as follows. The contributions to  $C_{lm}$  and  $S_{lm}$  values from all the lines in the Cartwright & Tayler expansion are pre-computed and saved in data files for each calendar day at 10-minute intervals. During orbit processing, the software reads these data files, and interpolates the contributions to the acceleration evaluation epoch. Further details are available in Bonin 2005. The ocean tide model used in RL06 is the same as the one used in RL05.

### II.2.4 Atmosphere & Oceanic Variability

The non-tidal variability in the atmosphere and oceans is removed through using the AOD1B Release-06 product. It is based on analysis and forecast data of the operational high-resolution global numerical weather prediction (NWP) model from the European Centre for Medium-Range Weather Forecasts (ECMWF) and ocean bottom pressure from an unconstrained simulation with a global ocean general circulation model that is consistently forced with ECMWF atmospheric data. The details of this product and its generation are given in the AOD1B Description Document (GRACE 327-750).  
[ftp://podaac-ftp.jpl.nasa.gov/allData/grace/docs/AOD1B\\_PDD\\_RL06\\_v6.0.pdf](ftp://podaac-ftp.jpl.nasa.gov/allData/grace/docs/AOD1B_PDD_RL06_v6.0.pdf)

This model of the geopotential is available as 3 hourly time series to degree and order 180. The value of the harmonics at intermediate epochs is obtained by interpolation between the bracketing data points.

In order to improve the accuracy of interpolation, the following procedure is adopted. First, the 3-hr epoch values from the AOD1B product are interpolated to 10-minute intervals called the “AOT” product. The Ray/Ponte  $S_2$  air tide model is also evaluated at 10-minute intervals and added back to the “AOT” product because the AOD1B RL06 does not include the  $S_2$  tide.

Just as with the ocean tides, the software reads these data files during orbit integration and interpolates the contributions to the acceleration-evaluation epoch.

### II.2.5 Solid Earth Pole Tide (Rotational Deformation)

The rotational deformation forces are computed as additions to spherical harmonic coefficients  $\bar{C}_{21}$  and  $\bar{S}_{21}$ , from an elastic Earth model, as specified in Section 6.4, IERS-2010.

Model	Description	Notes
An-Elastic Earth Model Contribution to C21 & S21	Scaled difference between epoch pole position and mean pole. See Chapter III (Earth Kinematics) for the linear	

	variation model for the mean pole.	
Polar Motion	Tabular input	<i>IERS C04</i>
Mean Polar Motion	Linear model	<i>IERS-2010</i> <sup>1</sup>

**II.2.6 Ocean Pole Tide**

The self-consistent equilibrium model of **Desai** is used (*Desai 2002*). A spherical harmonic expansion to degree 180 is used, with the same polar motion time series as for Earth Kinematics or the Solid Earth Pole Tide (See Section II.2.5).

The contributions to the spherical harmonic coefficients are pre-computed using software provided ([http://tai.bipm.org/iers/convupdt/convupdt\\_c6.html](http://tai.bipm.org/iers/convupdt/convupdt_c6.html)), and stored at 10-minute intervals for each calendar day. The orbit processing software reads these data files and interpolates the contributions to the integration or the acceleration evaluation epoch.

**II.2.7 N-Body Perturbations**

Unlike the geopotential accelerations, the perturbations due to the Sun, Moon and all the planets are directly computed as accelerations acting on the spacecraft. The direct effects of the objects on the satellite are evaluated using point-mass attraction formulas. The indirect effects due to the acceleration of the Earth by the planets are also modeled as point-mass interactions. However, for the Sun and Moon, the indirect effects include, in addition, the interaction between a point-mass perturbing object and an oblate Earth – the so-called Indirect J2 effect.

Model	Description	Notes
Third-Body Perturbation	Direct & Indirect terms of point-mass 3 <sup>rd</sup> body perturbations	
Indirect J2 Effect	Sun & Moon only	
Planetary Ephemerides	DE-430	

**II.2.8 General Relativistic Perturbations**

The general relativistic contributions to the accelerations and light propagation are computed as specified in Chapters 10 and 11 of the *IERS-2010 Conventions*.

**II. 3 NON-GRAVITATIONAL ACCELERATIONS**

The nominal approach is to use the GRACE-FO accelerometer data to model the non-gravitational accelerations acting on the satellite. The model used is:

$$\vec{f}_{ng} = q \otimes \left[ \vec{b} + {}_{3 \times 3} E \vec{f}_{acc} \right]$$

where the q/operator represents rotations to inertial frame using the GRACE-FO Attitude Quaternion SCA1B product; b represents an empirical bias vector; and the 3x3 matrix E

<sup>1</sup> See update at [http://maia.usno.navy.mil/conventions/2010/2010\\_update/chapter7/](http://maia.usno.navy.mil/conventions/2010/2010_update/chapter7/)

contains the scale factors along the diagonal and the cross-coupling or alignment factors in the off-diagonals. The bias vector and E matrix operate on the GRACE-FO Accelerometer observation ACC1B product, and are optionally estimable parameters.

**II. 4 EMPIRICAL ACCELERATIONS**

No empirical accelerations, either mean or once-per-revolution, were estimated.

**II. 5 NUMERICAL INTEGRATION**

The Predictor-Corrector formulation of the Krogh-Shampine-Gordon, second order, fixed-step, fixed-mesh/order integrator is implemented.

<b>Model</b>	<b>Description</b>	<b>Notes</b>
Dependent Variables	1. Equations of motion (position/velocity for each satellite) 2. State Transition Matrix (position/velocity mapping terms only)	
Formulation	Cowell Formulation	
Step-size and Order	5 seconds and 7 <sup>th</sup> order	

### III EARTH & SATELLITE KINEMATICS

#### III.1 EARTH ORIENTATION

Earth Orientation here refers to the model for the orientation of the Earth-fixed reference relative to the Inertial reference. The former are necessary for associating observations, models and observatories to the geographic locations; and the latter for dynamics, integration & ephemerides.

Frame	System	Realization
Inertial	ICRS	J2000.0 ( <i>IERS-2010</i> )
Earth-fixed	CTRS	IGS2014

The rotation between the Inertial and Earth-fixed frames is implemented as:

$${}_{3 \times 3} M_{trs}^{CRS} = Q(t)R(t)W(t)$$

which converts the column array of components of a vector in the terrestrial frame to a column array of its components in the inertial frame. Each component matrix is itself a 3x3 matrix, and is now individually described.

##### III.1.1 Precession & Nutation

Precession and Nutation are modeled using IAU2000A model (*Capitaine et al. 2002, Mathews et al. 2002*). Reference epoch 2000.0 is used. The independent variable is TT since epoch J2000.0 (noon, 01-Jan-2000).

##### III.1.2 Sidereal Rotation

This rotation is implemented as

$$R = R_3(-GST)$$

where the Greenwich Apparent Sidereal Time (GST) is calculated as the sum of Greenwich Mean Sidereal Time (GMST) and equatorial components of precession and nutation, consistent with the IAU 2000 resolutions. The GMST calculation uses UT1 as its independent argument, whose evaluation is also summarized in this table.

Quantity	Model	Notes
GMST	Polynomial with UT1 as independent variable.	<i>IERS-2003 Standards</i>
Equatorial components of precession & nutation	( <i>Aoki &amp; Kinoshita 1983</i> )	<i>IERS-2003 Standards</i>
UT1	cubic interpolation of tabular UT1 corrections	<i>IERS C04</i>
	Diurnal tidal variations from <i>Ray et al., 1994</i> eight constituent model.	<i>IERS-2010 Standards</i>

### III.1.3 Polar Motion

The Polar Motion component of rotation is implemented as

$$W = R_1(y_p)R_2(x_p)$$

Quantity	Model	Notes
Tabular variations	Cubic interpolation	<i>IERS C04</i>
Ocean Tidal Variations (Diurnal/Semi-Diurnal)	Orthoweights Formulation	<i>IERS-2010 Standards</i>
<p>Note 1: The rotation matrices are implemented in the small angle, skew-symmetric matrix formulation.</p> <p>Note 2: Rotational deformation accelerations &amp; kinematic station displacements are proportional to the difference between this time-series and a cubic mena-pole model.</p>		

### III. 2 STATION COORDINATES

This section summarizes the models for the mean and time-variable parts of the station coordinates adopted for data processing.

Quantity	Model	Notes
Mean Station Positions and Velocities	IGS14 (epoch-specific values are used)	Refers to the position of a geodetic marker or instrument reference point at each site.
Ocean Tidal Loading (Diurnal/Semi-Diurnal band)	Tidal orthoweights adjusted to site dependent displacement coefficients from Scherneck's loading service site.	<i>IERS-2010</i> Table 7.4, using GOT4.7 ocean tide model
Station Eccentricities	Individual observation models	
Luni-Solar Solid Earth Tidal displacement	Chapter 7, <i>IERS-2010</i> (Luni-Solar ephemerides from DE-430)	
Rotational Deformation	Scaling of difference of polar motion values from a cubic trend model.	Will switch to linear mean pole when a new IGS14 realization is produced consistent with that model.
Tidal Geocenter (Diurnal/Semi-Diurnal)	Included within Ocean Tidal Loading model	
Atmospheric Loading	Not modeled	No conventional model
Post-glacial Rebound	Not modeled	No conventional model
Seasonal Geocenter Variations	Not modeled	No conventional model

### III. 3 SATELLITE KINEMATICS

The Science Reference Frame (SRF - see *Product Specification Document*) is used in all instances where a satellite-fixed reference frame is needed. The inertial orientation of the spacecraft (i.e. the SRF) is modeled using tabular input quaternion from product SCA1B (*ibid.*). The same product is also used for rotating the accelerometer data to inertial frame prior to numerical integration; for making corrections to the ranging observations due to offset between the satellite center of mass & the antenna location; as well as for computing the non-gravitational forces (if necessary).

At epochs where the GRACE –FO quaternion product is not available, linear interpolation between adjacent values is used.

#### ***III.3.1 Rotation of Velocity Components***

The position rotations are specified in Section III. 1. The components of the satellite velocity vector are rotated using the matrix approximation

$$\vec{v}_{crs} = M_{crs}^{trs} \vec{v}_{trs} + (PNRS) \vec{r}_{trs}$$

#### ***III.3.2 GRACE-FO GPS Antenna and Laser Retroreflector Offset Model***

The GRACE-FO GPS navigation receiver is placed on the top surface (see *Product Specification Document*). For the purposes of orbit and gravity field determination, the antenna phase center location vectors for the L3 (LC) ionosphere-free double difference are

$$\begin{aligned} \text{GRACE-FO-1: } & (0.2602, -0.0013, -0.4770) \text{ meters} \\ \text{GRACE-FO-2: } & (0.2600, -0.0011, -0.4762) \text{ meters} \end{aligned}$$

in the Science Reference Frame. This is consistent with the value provided in the VGN1B product (RL04). The IGS14 antenna phase-center variations (PCV) maps for the GPS transmitter satellites and for the ground-stations were modeled, as were the GRACE-FO receiver antenna PCV maps.

The GRACE-FO laser retroreflector array (LRA) coordinates are

$$(-0.6000 \ -0.3275 \ 0.2188) \text{ meters}$$

The LRA design is identical to GRACE, so that an additional 4 mm correction is required to model the localization of the LRA reference point. This correction actually varies according to the angle of incidence with a range of approximately  $\pm 2$  mm (Grunwaldt, 2001).

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